



ESA-MOST China Dragon 4 Cooperation

→ ADVANCED TRAINING COURSE IN OCEAN AND COASTAL REMOTE SENSING

12 to 17 November 2018 | Shenzhen University | P.R. China

Application of altimetry in the China Seas Jungang Yang (杨俊钢), First Institute of Oceanography, State Oceanic Administration





outline

- 1. Ocean tide
- 2. Ocean wave
- 3. Sea level change
- 4. Ocean circulation and mesoscale eddy
- 5. Marine gravity anomaly







What can do?

Introduce the main applications of altimeter data in the ocean research.

How to do?

Introduce the basic method or technical flow for the applications of altimeter data.

What have done?

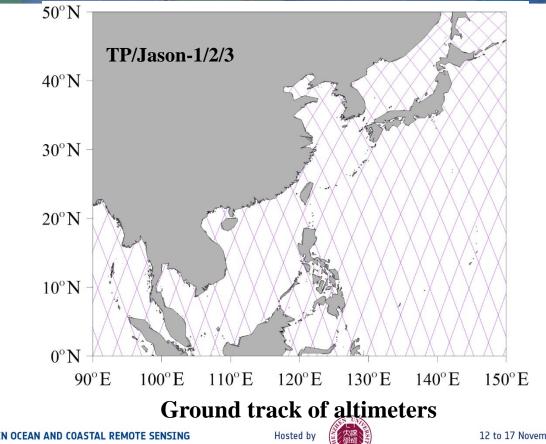
Give a review about the altimeter application in the China Sea.





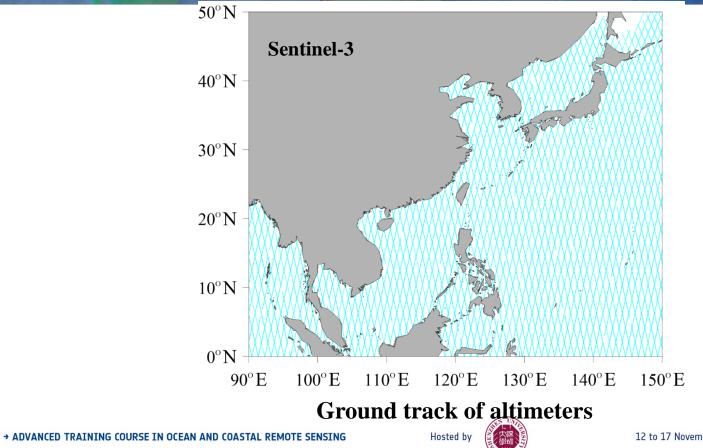
Satellite	Country/Organization	Orbit/km	inclination/°	Repeat Cycle/d	Period
T/P	USA/France	1336	66	9.9156	1992.9.25~2005.10.8
ERS-1	ESA	799	98.55	35/30	1991.8.1~1996.6.2
ERS-2	ESA	799	98.55	35/30	1995.4.29~2010.9.13
GFO	USA	880	108	17	2000.1.9~2008.9.7
Jason-1	USA/France	1336	66	9.9156	2002.1.15~2012.12.11
Envisat	ESA	799	98.55	35/30	2002.5.14~2012.4.8
Jason-2	USA/France	1336	66	9.9156	2008.7.4~
Cryosat-2	ESA	717	92	369	2010.7.16~
HY-2A	CHINA	973	98	14	2011.10.01~
SARAL	India/France	800	98.55	35	2013.3.14~
Jason-3	USA/France	1336	66	9.9156	2016.2.12~
Sentinel-3	ESA	814.5	98.65	27	2016.12.24~
HY-2B	CHINA	973	98	14	2018.10.25~



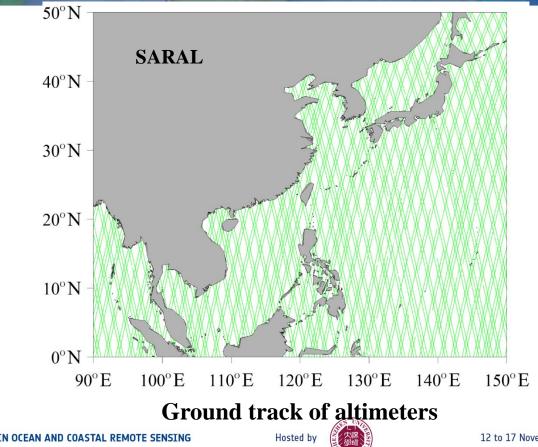


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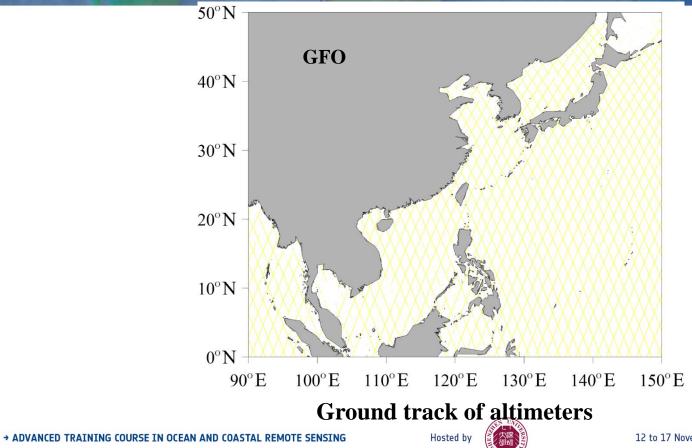




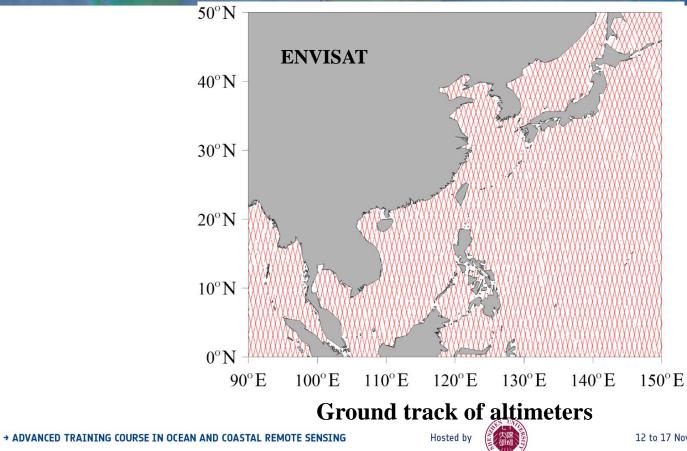




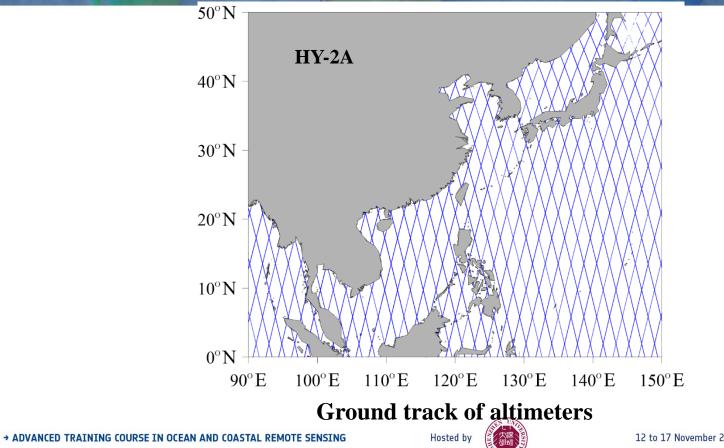




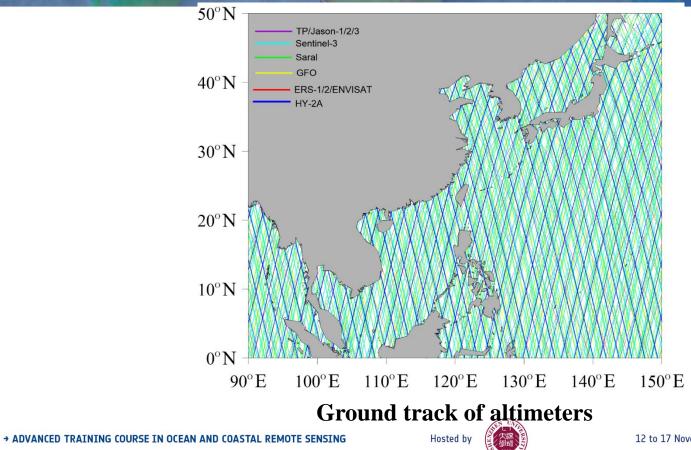








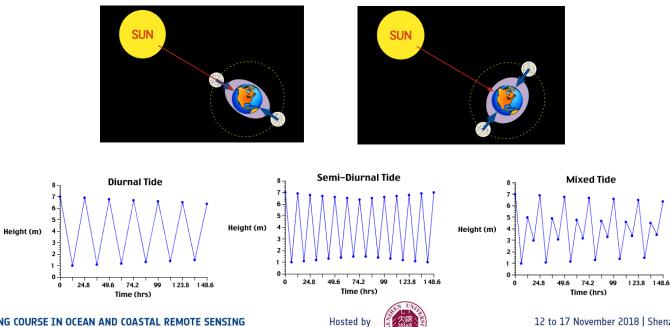




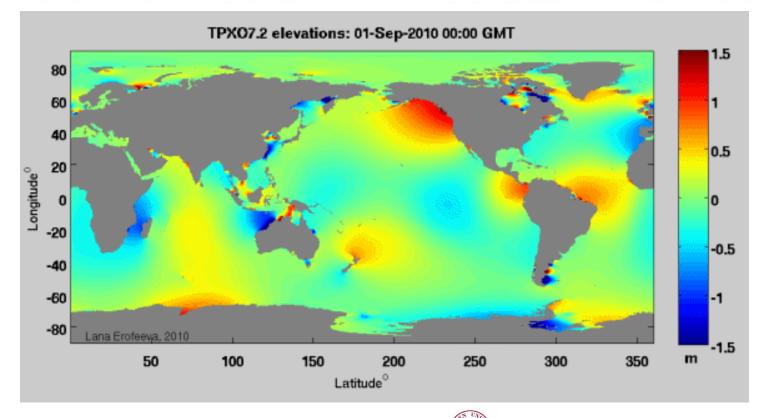
1. Ocean Tide



Ocean tide is of importance in the study on oceanography, geophysics and geodesy.





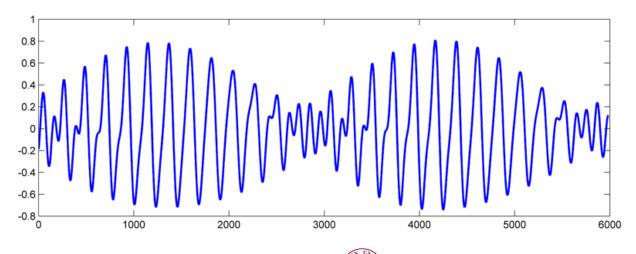






(1) Extraction of tide with altimeter data

Based on the time series of altimeter data, the harmonic constants of ocean tide can be extracted by two methods.







➤ Harmonic Analysis Method(调和分析法)

Ocean tide elevation can be regarded as the sum of several tidal constituents of different frequencies.

$$\xi(\vec{x},t) = \sum_{i=1}^{N} f_i(t) A_i(\vec{x}) \cos[\omega_i t + V_i + u_i(t) - G_i(\vec{x})]$$

Where, $f_i(t)$, $u_i(t)$ are the nodal corrections of tidal constituents.

 $A_i(\vec{x}), G_i(\vec{x})$ are the amplitude and phase lag of tidal constituents. V_i is the astronomical phase, and ω_i is the frequency of the tidal constituents.





> Orthogonalized Convolution Method(正交响应法) Ocean tide elevation is expressed as the sum of a series of orthotides.

$$\xi(t) = \sum_{n=2}^{\infty} \sum_{m=0}^{n} \sum_{k=-K}^{K} [u_n^m(k)a_n^m(t-k\Delta t) + v_n^m(k)b_n^m(t-k\Delta t)]$$

The orthotides expression of Groves and Reynolds(1990): $\xi(t) = \sum_{m=1}^{2} \sum_{j=0}^{2} [U_{j}^{m} P_{j}^{m}(t) + V_{j}^{m} Q_{j}^{m}(t)]$

 U_i and V_i can be obtained by least square method.

 $H\cos g = H_1 = (-1)^m H_0 X(\omega)$ $H\sin g = H_2 = (-1)^m H_0 Y(\omega)$

 $X(\omega)$, $Y(\omega)$ can be calculated by U_j and V_j . H_0 is the amplitude of astronomical equilibrium tide potential. *H* and *g* are the harmonic constants of tidal constituents.



(2) Combination of altimeter data with numerical model

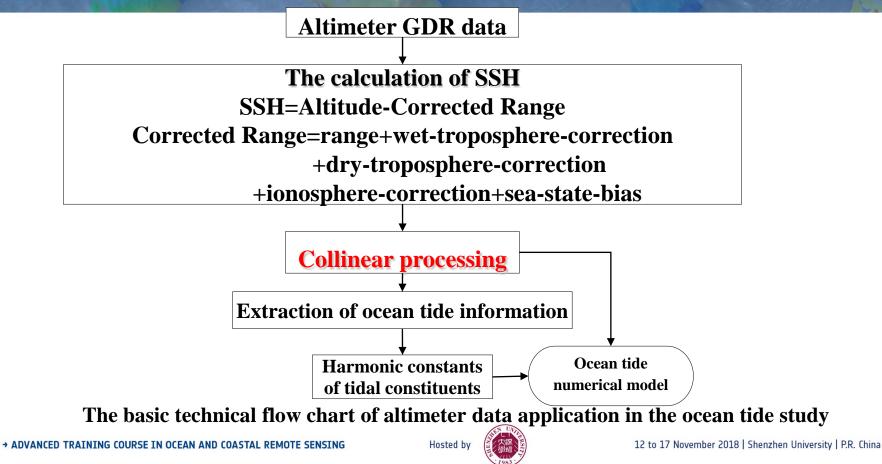
> example: 2-D ocean tide numerical model

$$\begin{cases} \frac{\partial \zeta}{\partial t} + \frac{\partial (Hu)}{\partial x} + \frac{\partial (Hv)}{\partial y} = 0\\ \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} - fv = -g \frac{\partial \zeta}{\partial x} - \frac{C_b u}{H}\\ \frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + fu = -g \frac{\partial \zeta}{\partial y} - \frac{C_b v}{H} \end{cases}$$

Where, u, v are the east and north components of tidal current. ζ is tidal elevation. f is Corolis parameter. C_b is the bottom friction coefficient. g is the gravity acceleration. H is water depth.

solving the above numerical model, altimeter data can be used as the input of water elevation ζ in the open boundary. Advanced TRAINING COURSE IN OCEAN AND COASTAL REMOTE SENSING







Collinear processing: using interpolation method to process the altimeter data of each cycle into a reference ground track with the fixed ground points. That makes to obtain the time series of SSH observations at the same ground track point.

Cycle *i* Cycle *j*

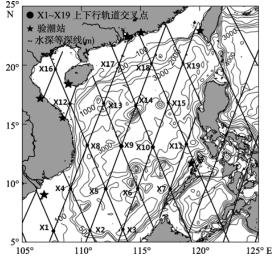
Processing the same pass in different cycles

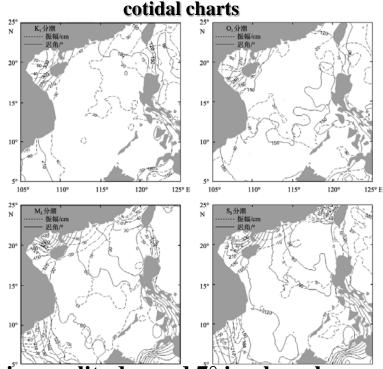


1.2 The application in the China Seas



10 years (1992~2002) of T/P data is performed to derive the semidiurnal and diurnal tides in the South China Sea.





Comparisons with TPXO7.2 is less than 3 cm in amplitudes and 7th in phase lags.¹²⁵

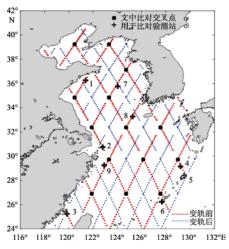
Yunxia ZHAO, et. al., 2012, Marine Sciences, Vol.36(5).(In Chinese)

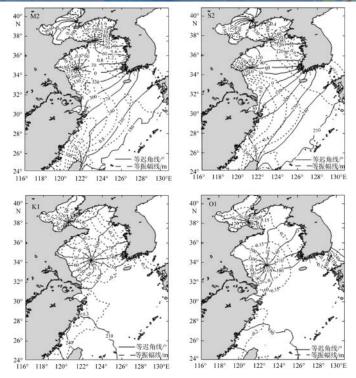
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> The harmonic analysis of 8 tidal constituents in the East China Sea, Yellow Sea, Bohai Sea was performed by combining of 19 years altimeter data of T/P and Jason-1.





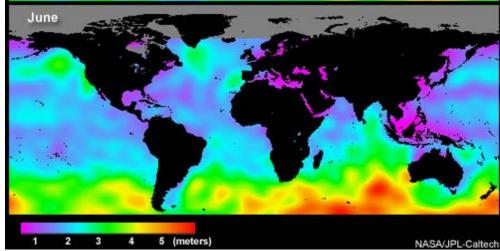
The RMSE of Amplitude is 1.0~1.8 cm and that of phase lag is 4.1° ~7.8°,

Changwei ZHONG, et. al., 2013, Marine Sciences, Vol.37(10).(In Chinese) Hosted by → ADVANCED TRAINING COURSE IN OCEAN AND COASTAL REMOTE SENSING



Ocean wave is one of the basic parameter of sea state, and it reflects the local ocean and weather condition.









Wave/wind observation: obtain SWH, sea surface wind speed and monitor sea state. Statistic characters analysis of ocean wave.

Extreme SWH estimation: extreme SWH in many return years can be estimated by using altimeter data.

Model validation: validate the result of numerical simulation and the statistic distribution model of ocean wave.

Ocean wave data assimilation: assimilate SWH data into the numerical ocean wave model (WAM, SWAN or WAVEWACTH III) by the optimal interpolation or other methods.

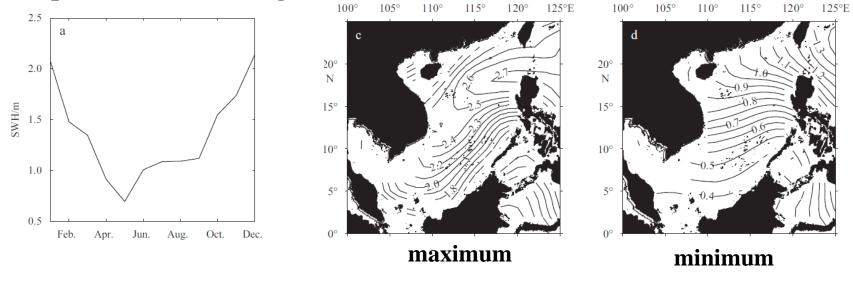


SARSEE CESA

2.2 The Application in the China Seas



The seasonal variability of the SWH in the South China Sea is investigated using the gridded daily altimeter data for the period of September 2009 to August 2015.



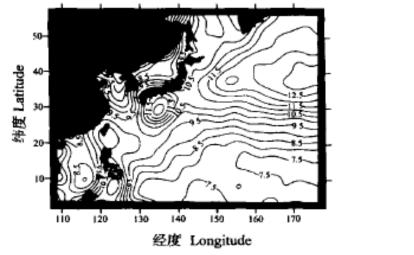
SU Hui et al. Acta Oceanol. Sin., 2017, Vol. 36, No. 11, P. 38-50.

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The estimation of extreme SWH using T/P data, the results were compared with those calculated by traditional method and observed data.



经度 Longitude 经度 Longitude 经度 Longitude Threshold of SWH (100 Return Years) Threshold of SWH (50 Return Years)

Han Shuzong, 2003. Journal of ocean university of Qingdao, 33(5), 657-664.

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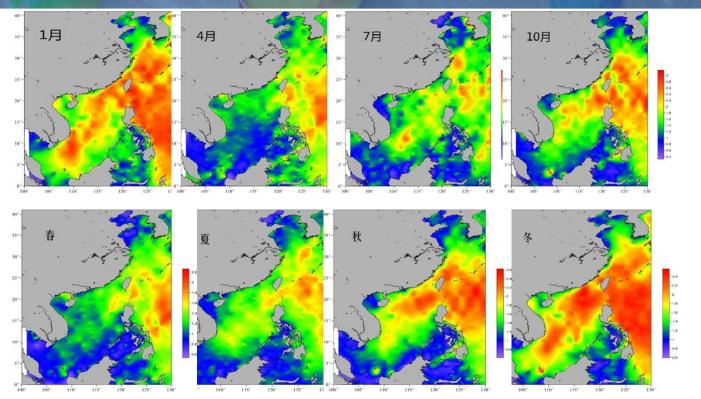
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Latitude

520-





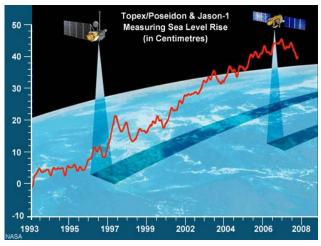
Monthly and seasonal mean distribution of SWH in the China Sea based on Jason-1 data



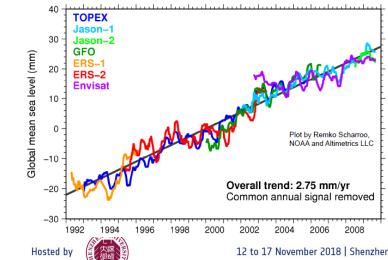
3. Sea Level Change



- The study on sea level change is the important section of the global change study. This study aims to understand the variation of sea level, to forecast the future and its impacts on the environment of human living.
- Altimeter supplies the unique data for the study on the global sea level change, which make up the shortcoming of tide gauge data.









Altimeter data can be used in the study on:

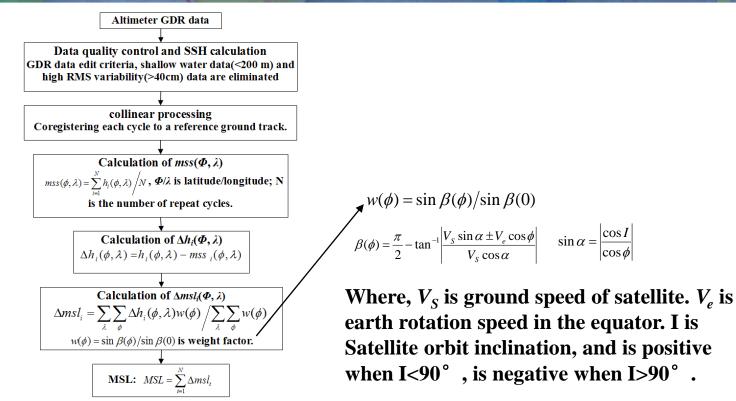
- The trend of sea level rise.
- The sea level rise rate.
- Analyzing the main reasons of sea level change in combination with other data, such as SST, tide gauge, glacier or ice sheet data.



NRSCC

esa





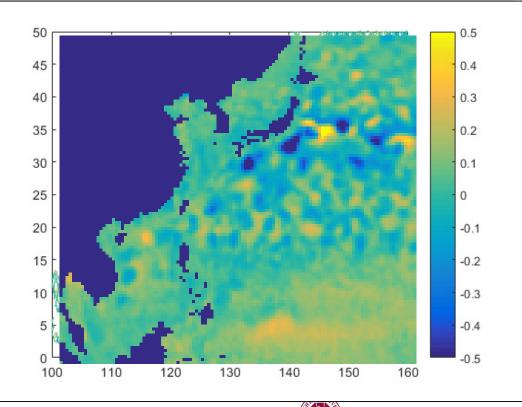
The technical flow chart of altimeter data application in the sea level change study

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Along track SLA ----→ gridding SLA

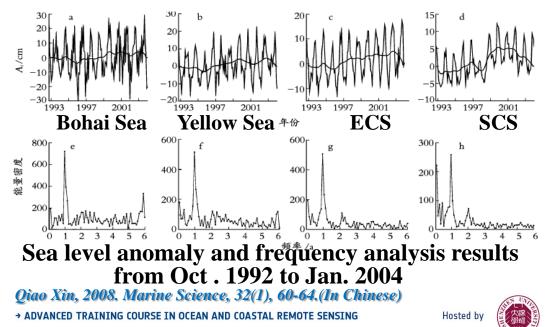


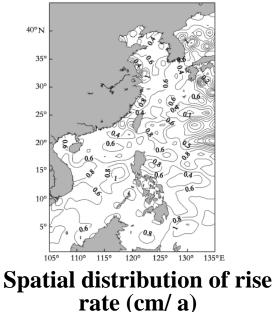




3.2 The application in the China Seas

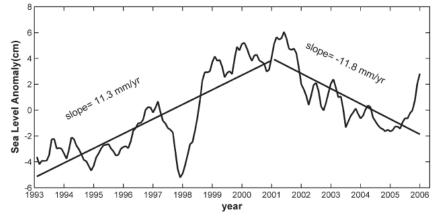
- Surse Cesa
- The period characteristics and rise rate of the sea level in the China Sea were analyzed based on T/P and Jason-1 data from 1992 to 2004, the rise rate was 0.593cm/a.







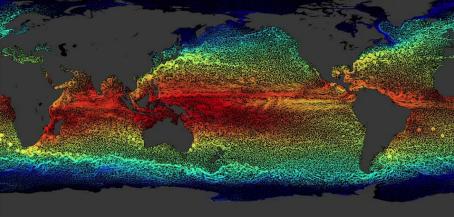
Sea surface variability in the South China Sea was analyzed to show that the rise rate and fall rate from 1993 to 2006 were 11.3 mm/yr and 11.8 mm/yr respectively.



Time evolution of the SCS mean sea level anomaly from January 1993 to December 2005

4. Ocean circulation and mesoscale eddy

- Ocean circulation and mesoscale eddy are the important forms of sea water movement, and they are important means of transport of mass and energy in the ocean.
- Based on altimeter data, the global ocean circulation and mesoscale eddy were understood.



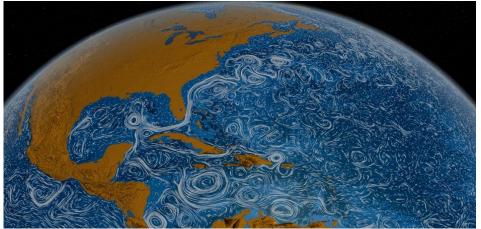


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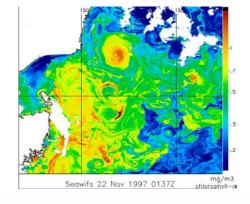
esa

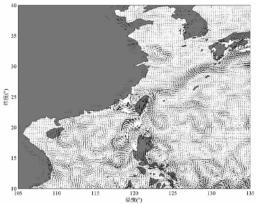
Mesoscale Eddy

- Large bodies of swirling water
- Nearly everywhere in the global oceans
- **Temporal scale:** ten days ~ several months \geq
- Spatial scale: ~100 km \succ













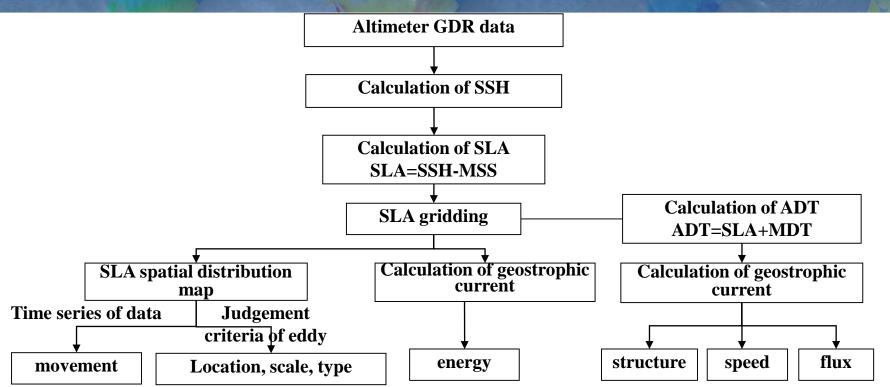


4.1 Application in the ocean circulation and mesoscale eddy study

With altimeter data, we can study:

- Structure, speed and flux of ocean circulation.
- Annual variability of ocean current.
- Characters, movement and energy of mesoscale eddy.
- Data assimilation to the numerical model.





The technical flow chart of data application in the ocean circulation and mesoscale eddy study

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According to geostrophic balance, the pressure gradient force is equal to Corolis force.

$$\begin{cases} u = -\frac{g}{f} \frac{\partial h}{\partial y} \\ v = \frac{g}{f} \frac{\partial h}{\partial x} \end{cases}$$

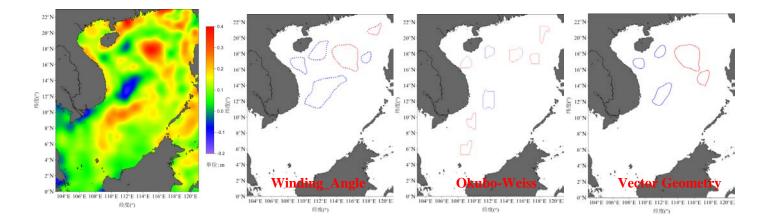
Where, u, v is the east and north component of geostrophic current. f is Corolis parameter. g is gravitational acceleration. h is ADT.



Eddy detection Method



> Winding Angle, Okubo-Weiss parameter, Vector Geometry



The Winding Angle presents the best performance, similar to the results by other scientists (Chaigneau et al., 2010; Souza et al., 2011).





Winding Angle method

edge.

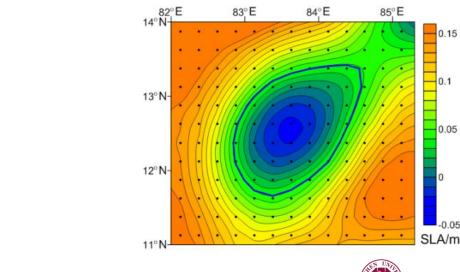
- to identify the local SLA maximum or minimum as the center of eddy.
- **(2)** to select outermost closed SLA contour associated to each eddy as the eddy

0.15

0.1

0.05

0.05





The eddy radius *R* is defined as the equivalent radius of a circle with the same area *A* as the closed contour of the eddy borders,

$$R = \sqrt{A/\pi}$$
.

The amplitude is defined as the SLA difference between the center and bourder of mesoscale eddy,

$$Amplitude = |SLA_{max} - SLA_{border}|.$$

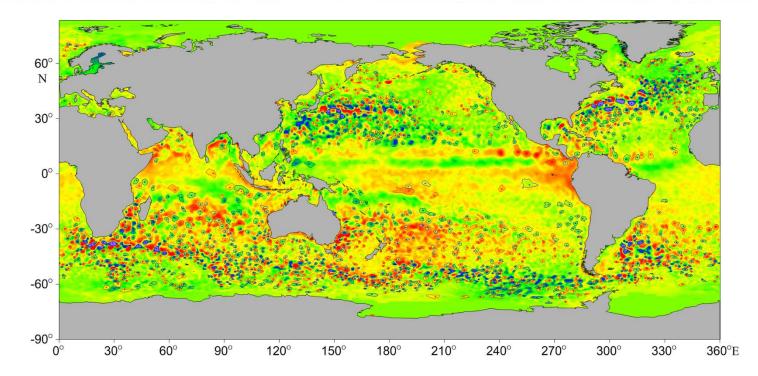
The eddy intensity is characterized by the eddy kinetic energy EKE,

$$EKE = \frac{1}{2} (u'^2 + v'^2).$$



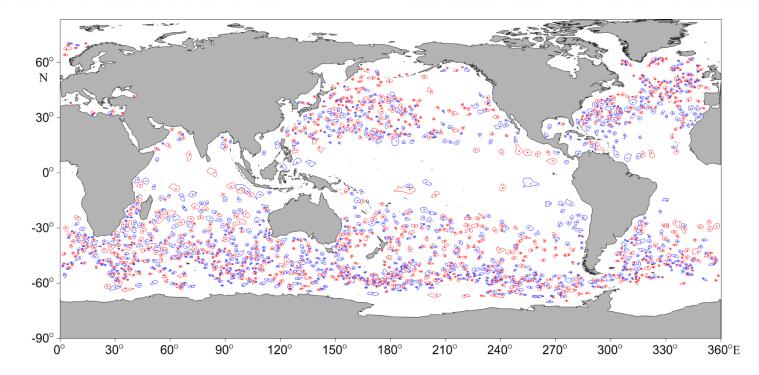
Mesoscale eddies detection results





SLA on May 3, 2011

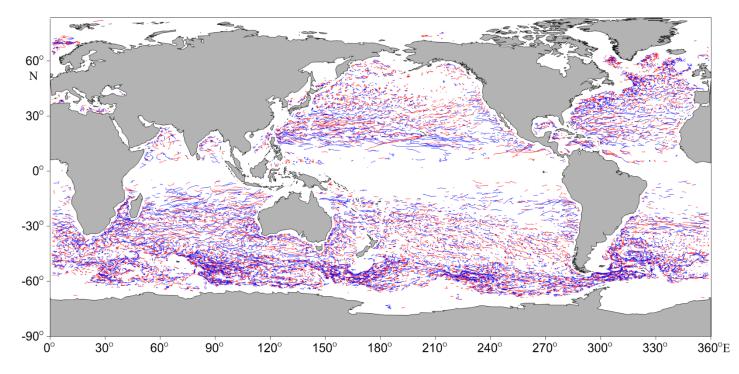




Eddy detection result on May 3, 2011

Eddy Tracking result

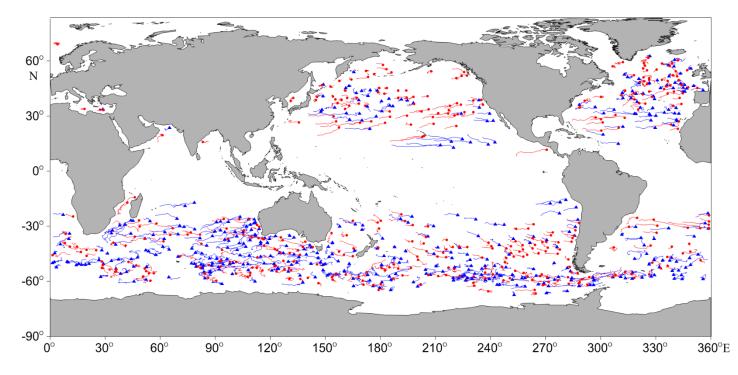




The propagation trajectories of cyclonic and anticyclonic eddies > 30d on 2011

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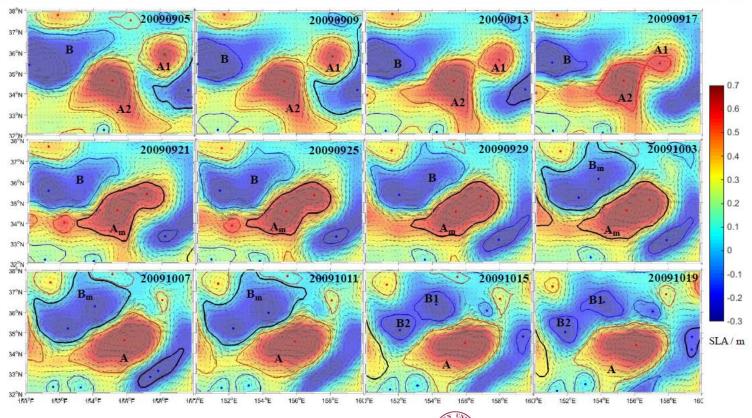




The propagation trajectories of cyclonic and anticyclonic eddies >180d on 2011

Multicore eddy evolution

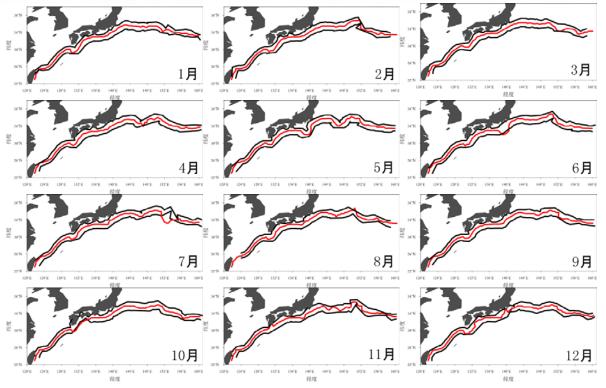






4.2 The application in the China Seas

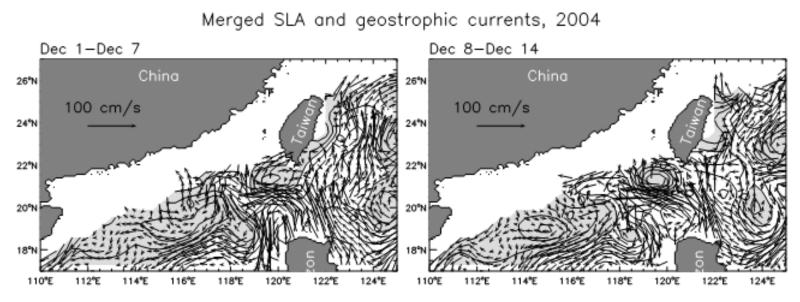




Annual average of the Kuroshio



altimeter data, combining with Satellite ocean color and SST data, are used to study the surface Kuroshio path in the Luzon Strait area.



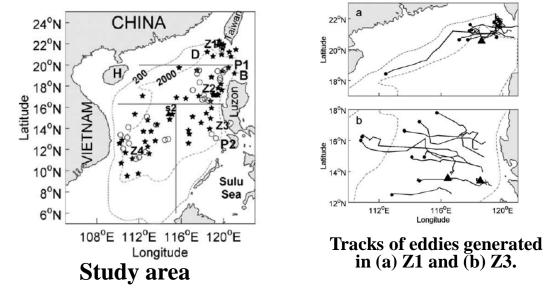
Dongliang Yuan, 2006. JGR, 111, C11007, doi:10.1029/2005JC003412.

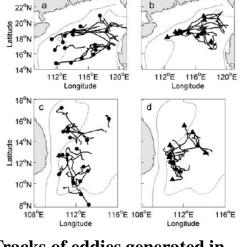
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A composite time series (1993–2000) of SSHA from several satellites is used to identify eddies in the SCS. The eddy lifetime, radius, strength, and straight-line travel distance are estimated.





Tracks of eddies generated in Z1 (a, b) and Z4 (c, d).

Guihua Wang et al., 2003. GRL, 30(21), doi:10.1029/2003GL018532.

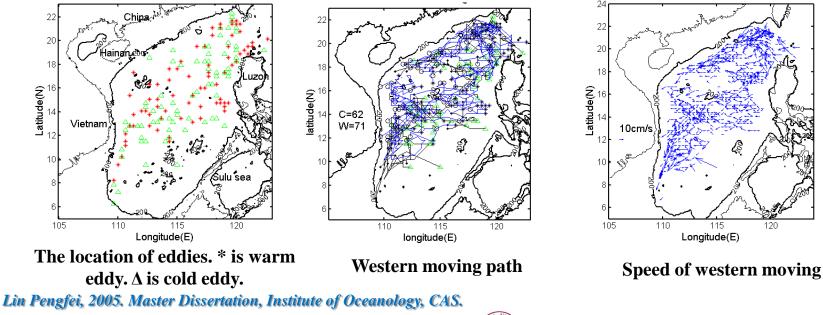
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120°E



Based on T/P and ERS-1/2 merged data, the mesoscale eddies in the SCS from 1993 to 2001 were identified and tracked, and theirs characters of generation and movement were analyzed.

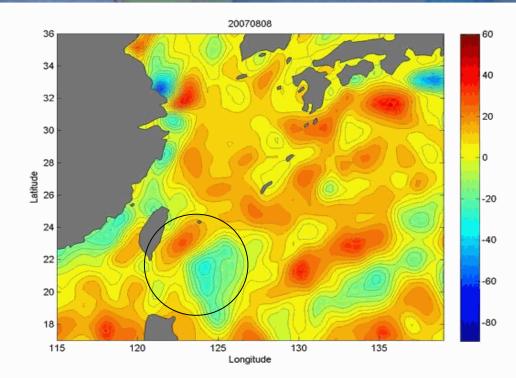




120

115

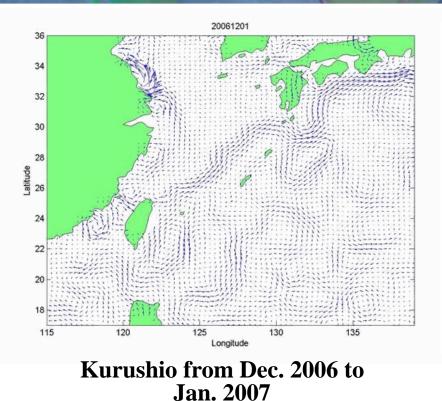


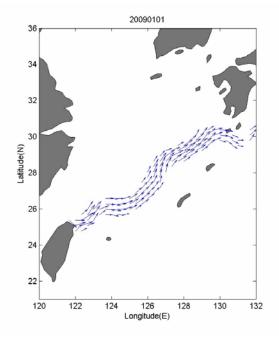


The evolvement of a cold eddy in the eastern Luzon strait during 2007.8.8-2007.10.10









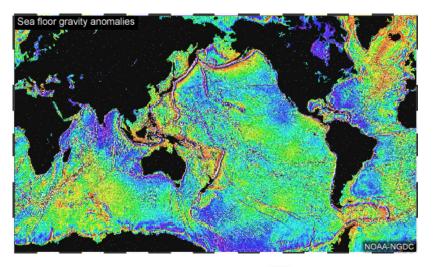
The main current variability of Kuroshio in 2009



5. Marine Gravity Anomaly



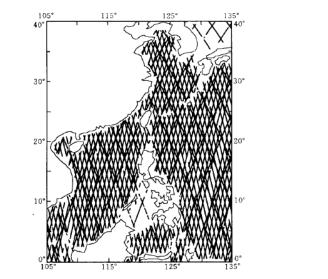
- The determination of the marine gravity field is very important for studying the global gravity field.
- Satellite altimetry has been the important mean to obtain the information of marine gravity field now.

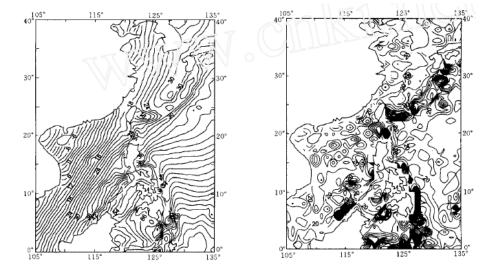






Based on T/P and ERS-1 data, the 30'×30' gravity anomalies in the China Seas and its vicinity are recovered. The precision was 3.5mGal.





Geoid undulations (left) and gravity anomalies in the region of China seas and vicinity

Hsu Hou-tse et al., 1999. Chinese Journal of Geophysics, 42(4), 465-471.

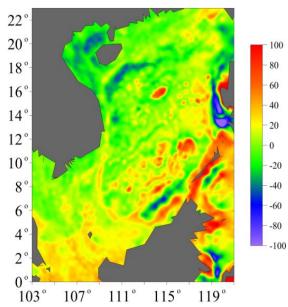
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Ground tracks of T/P and ERS-1





gravity anomalies inversion by multi-source altimeters data (Geosat, T/P, ERS-1/2, Envisat, Jason-1/2/3, HY-2A RA).



Compare to in-situ measurement data, the RMSE is less than 5.8 mGal.







Ocean Tide

- > extract harmonic constants of tidal constituents.
- > Assimilate altimeter data into the numerical ocean tide model.
- The application in the China Sea: extracting diurnal, semidiurnal and other tidal constituents based on altimeter data.







Ocean Wave

- > Analyze statistic characters of ocean wave.
- ➢ Estimate extreme SWH.
- > Validate the statistic distribution and numerical model.
- > Assimilate altimeter SWH data into the ocean wave model.







Sea Level Change

- Trend and rate of sea level rise.
- > Analyze the main reasons of sea level.
- > Altimeter data assimilation of the numerical model.







Ocean Circulation and Mesoscale Eddy

- > Structure, speed and flux of ocean circulation.
- > Annual variability of ocean current.
- > Characters, movement and energy of mesoscale eddy.
- Application in the China Seas: Kuroshio, surface circulation and mesoscale eddies in the SCS, ECS and Yellow sea, and the data assimilation of the numerical model.







Marine Gravity Anomaly

- > Recover the gravity anomaly with satellite altimeter.
- The gravity anomalies of different spatial resolution in the China seas were recovered by combining multi-altimeter data.





Thanks for your attention!

Email: yangjg@fio.org.cn

