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Qualitative landslide susceptibility assessment by multicriteria analysis: A case study from San Antonio del Sur, Guantánamo, Cuba

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Abstract

Geomorphological information can be combined with decision-support tools to assess landslide hazard and risk. A heuristic model 9 was applied to a rural municipality in eastern Cuba. The study is based on a terrain mapping units (TMU) map, generated at 1:50,000 10 scale by interpretation of aerial photos, satellite images and field data. Information describing 603 terrain units was collected in a 11 database. Landslide areas were mapped in detail to classify the different failure types and parts. Three major landslide regions are 12 recognized in the study area: coastal hills with rockfalls, shallow debris flows and old rotational rockslides denudational slopes in 13 limestone, with very large deep-seated rockslides related to tectonic activity and the Sierra de Caujerí scarp, with large rockslides. The 1415Caujerí scarp presents the highest hazard, with recent landslides and various signs of active processes. The different landforms and the causative factors for landslides were analyzed and used to develop the heuristic model. The model is based on weights assigned by 16expert judgment and organized in a number of components such as slope angle, internal relief, slope shape, geological formation, 17 active faults, distance to drainage, distance to springs, geomorphological subunits and existing landslide zones. From these variables a 18 hierarchical heuristic model was applied in which three levels of weights were designed for classes, variables, and criteria. The model 19combines all weights into a single hazard value for each pixel of the landslide hazard map. The hazard map was then divided by two 20 scales, one with three classes for disaster managers and one with 10 detailed hazard classes for technical staff. The range of weight 2122 values and the number of existing landslides is registered for each class. The resulting increasing landslide density with higher hazard classes indicates that the output map is reliable. The landslide hazard map was used in combination with existing information on 23buildings and infrastructure to prepare a qualitative risk map. The complete lack of historical landslide information and geotechnical 24data precludes the development of quantitative deterministic or probabilistic models. 25

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Keywords: Landslide susceptibility; Geomorphological mapping; Heuristic analysis; Multicriteria analysis

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1. Introduction

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As in most developing countries, landslide inventory 31 maps are still scarce in Cuba, due to the limited resources 32 available for research. Most conventional landslide 33 studies in Cuba are descriptive, and a few focus on 34 hazard assessment. Moreover, most of the quantitative 35

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risk assessment methods that have been developed
elsewhere are case-specific and require many types of
data, on landslide occurrence and impact, most of which
are not yet available in Cuba.

There are four different approaches to the assessment 40 of landslide hazard: landslide inventory-based probabi-41 listic, heuristic (which can be direct geomorphological 42mapping, or indirect qualitative map combination). sta-43 tistical (bivariate or multivariate statistics) and deter-44 ministic (Soeters and Van Westen, 1996; Aleotti and 45 Chowdhury, 1999; Guzzetti et al., 1999). Landslide risk-46 assessment methods are classified into three groups, as 47 qualitative (probability and losses in qualitative terms), 48 semi-quantitative (indicative probability, qualitative 49terms) and quantitative (probability and losses both nu-50merical) (Lee and Jones, 2004). 51

The heuristic approach is considered to be useful for 52obtaining qualitative landslide hazard maps for large 53areas in a relatively short time. It does not require the 54collection of geotechnical data, although detailed geo-55morphological mapping is essential. The heuristic ap-56proach may result in more reliable susceptibility maps 57than using statistical methods, where a considerable 58 59amount of generalization always needs to be accepted in the analysis. This is particularly relevant in Cuba, where 60 the geomorphological setting exhibits high spatial var-61 iability. Qualitative risk assessment procedures in many 62 countries are heuristic, e.g. in Switzerland (Lateltin, 63 1997). The qualitative approach is based on expert opin-64 ion and the risk areas are categorized by such terms as 65 "very high", "high", "moderate", "low" and "very low". 66 The increasing popularity of geographic information 67 systems (GIS) has lead to many studies, mainly using 68 indirect susceptibility-mapping approaches (Aleotti and 69 Chowdhury, 1999). As a consequence fewer investiga-70 tions use GIS in combination with a heuristic approach, 71 either geomorphological mapping, or index overlay 72mapping (e.g. Barredo et al., 2000; Van Westen et al., 73 2000, 2003). 74

Nowadays, new decision-support tools are available 75 for GIS-based heuristic analysis. They allow better 76 structuring of various components, including both ob-77 jective and subjective aspects and compare them in a 78 logical and thorough way (Saaty, 1980). Decision-79 support tools such as (spatial) multicriteria analysis have 80 not been popular for qualitative assessment of landslide 81 hazard. 82

This research combines geomorphological information with GIS-based decision-support tools to develop a heuristic landslide-hazard assessment model in San Antonio del Sur, Guantánamo, Cuba at 1:50,000 scale.

2. The study area

The study area, within the San Antonio del Sur 89 municipality, is located in eastern Cuba (Fig. 1) 60 km 90 from the city of Guantánamo, the capital of the province 91 with the same name. The main access to the area is by 92 the coastal road connecting Guantánamo and the eastern 93 municipalities. 94

The geology and tectonic setting of the eastern part of 95 Cuba is rather complicated, and includes several 96 geological and tectonic environments in a relatively 97 small area. The different tectonic and structural processes 98 have overlapped over geological time in such a way that it 99 is difficult to separate them spatially and temporally. 100 Moreover, the area remains an active tectonic zone on the 101 northern boundary between the Caribbean and North 102 American plates, as evidenced by many neotectonic 103 features and by the continuous general uplift of the area. 104 The general geology of Cuba is divided into two principal 105 units: a foldbelt and a neoautochthon (Iturralde-Vinent, 106 1996), which unconformably overlies the foldbelt 107 (Fig. 1). The eastern part of Cuba and the study area 108 have been studied by Nuñez et al. (1981), Nagy et al. 109 (1983), Millán and Somin (1985) and Franco (1992). 110

The San Antonio del Sur area contains geological units 111 from both the foldbelt and the neoautochthon. The foldbelt 112 consists of a Northern Ophiolite belt, a Cretaceous 113 Volcanic arc, a Paleocene Middle Eocene volcanic arc 114 and a Late Middle Eocene piggyback basin (Fig. 1). The 115 ophiolites are represented in the study area by a hilly zone 116 in the southeast called "Sierra del Convento" (Fig. 2). It is 117 the surface expression of a larger body considered as a 118 relic of the basement, which emerged due to spreading of 119 the oceanic crust, pushed from the south by tectonic events 120 (Chang and Suarez, 1998). Rocks of the Cretaceous 121 volcanic arc belonging to the Sierra del Purial Formation 122 underlie the eastern part of the study area with low grade 123 and high pressure metamorphism. Deposits from the 124 Paleocene volcanic arc, belonging to the El Cobre Group, 125 can be found in the north of the area (Fig. 1). The 126 piggyback basin corresponds to the Paleocene-middle 127 Eocene volcanic arc. The Charco Redondo, San Ignacio 128 and San Luis formations represent this basin in the study 129 area. These formations cover the northeast up to the central 130 part, and consist of polymictic sandstones, mudstones, 131 marls, clays, limestone clays, bioclastic limestone, sandy 132 limestone and polymictic conglomerates.

The western and central part of the study area is 134 underlain by the "Neoautochthon", recent units formed in 135 situ, and which are represented in the study area by 136 formations from three transgression–regression phases 137 which occurred since the Late Eocene (Iturralde-Vinent, 138

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Fig. 1. A: Location of study area in Cuba. 1 = outcrops of the foldbelt, 2 = Eocene to Recent neo-autochthonous deposits, 3 = Main faults, 4 = Study area. B: general geological map of the San Antonio del Sur municipality, Guatánamo province, Cuba. C: Shaded relief of Eastern Cuba with study area outlined. North is up and the area is 20 by 30 km.

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Fig. 2. Geomorphological map. San Antonio del Sur, Guantánamo Province, Cuba.

1996). The first such cycle is from Late Eocene to
Oligocene (Phase1) and consists of alternating layers of
sandstone, mudstone and calcareous clay. The second
cycle is from Early Oligocene–Late Miocene (Phase2),
and is characterized by alternating layers of clastic,
bioclastic and biogenic limestone (Fig. 1B). The third
cycle is Late Pliocene to recent (Phase3) and is

characterized by algae-bearing bioherm limestones, with 146 corals, and recent marine sediments.

3. Geomorphology and landslide occurrence 148

The geomorphology of the study area is conditioned 149 by the Caribbean–North American inter-Plate zone, and 150

the paleoclimatic oscillations during the Quaternary
period. Fig. 3 presents an overview of the study area as
an anaglyph image, generated from a digital elevation
model and a SPOT-PAN image. For 3-D viewing of the

image, red-green glasses are required. The inter-Plate 155 boundary consists of a strike-slip fault system (Fig. 1) 156 of which secondary faults have topographic effects 157 inland (see Fig. 3, complexes C and B). The Quaternary 158



Fig. 3. Geomorphic complexes of the study area: A, B, C: denudational hills in metamorphic rocks, terrigenous rocks, and in limestone respectively; D: Alluvial units; E: Accumulational slopes; F: Coastal hills; G: Puriales de Caujeri depression. This anaglyph image requires red–green glasses for viewing in 3-D.

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sea-level oscillations left coastal hills with marine
terraces and abandoned valleys (see Fig. 3, complexes
D, E and F).

A geomorphological map, including the landslide 162 inventory, (Fig. 2) at a scale of 1:50,000 was prepared 163 from interpretation of two sets of aerial photographs (of 164 1:25,000 and 1:37,000 scale) and fieldwork. Both photo 165sets correspond to a national aerial survey carried out at 166 the beginning of the 1970s. The 1:37,000 scale photos 167 (55 in total) cover the whole study area with four flight 168 lines and were taken between 2-Feb-1972 and 19-Mar-169 1972. The 1:25,000 scale photos (46 in total) cover the 170 171 south-west part of the study area in three flight lines and were taken between 5-Dec-1971 and 21-Dec-1971. The 172photos were interpreted with a TOPCON stereoscope on 173transparent paper and transferred to digital format by on-174screen digitizing using other image products for double 175checking (anaglyph, shaded DEM, Landsat TM true 176color composite and digital topographic map). The 177 photo-interpreted units were checked in the field by 178three people during a fieldwork campaign which took 179three weeks. 180

The area was divided into 603 terrain mapping units 181 182 (TMU). A TMU can be considered a homogeneous mapping unit on the basis of geomorphologic origin, 183 physiography, lithology, morphometry, and soil geogra-184 phy (Meijerink, 1988). Normally, a single landslide was 185 considered an individual TMU. In certain cases, when the 186 size was large enough, landslide zones such as scarps, 187 bodies and depressions were also considered as separate 188 TMUs. The units have been combined into the following 189 geomorphological complexes: 190

- Denudationalhills (metamorphic rocks, terrigenous rocks, and limestone) (A, B and C in Fig. 3)
- Alluvialunits and accumulational slopes (D and E in
 Fig. 3)
- 195 Coastal hills (F in Fig. 3)
- ¹⁹⁶ Purialesde Caujeri depression (G in Fig. 3)

197 3.1. Denudational hills

Denudational hills can be separated by the underly-198ing lithology into a limestone plateau (C in Fig. 3), 199 metamorphic hills (A in Fig. 3) and terrigenous hills (B 200in Fig. 3). The plateau is a monocline at an average 201 altitude of 500 m, composed mainly of limestone from 202 the Yateras formation (Phase2) with extensive karst 203processes and additional erosion where the underlying 204 mudstone and calcareous clay from the Maquey for-205mation (Phase3), are exposed. The area is dominated by 206neotectonic processes such as faulting, and a number of 207

large deep-seated mass movements are observed which 208 could reflect a mix of gravitational and tectonic move- 209 ments (Fig. 3). The landslides are concentrated in the 210 north part of Baitiquirí and El Naranjo. As the area is 211 strongly affected by active faults the landslides are 212 located in three main "steps" from the limestone hills 213 toward the coast. In the upper part the landslides occur 214 in limestone rocks of Yateras formation (Fig. 2). In the 215 next two steps the landslides occur over terrigenous 216 materials (Fig. 1). All the main landslide features are 217 aligned with the major faults in the directions SW-NE, 218 N-S and W-E, and they occur often where the faults are 219 converging. In some of the upper portions of the large 220 landslide masses the total amount of displacement has 221 not been more than a few meters. Displacements were 222 greater on the lower steps and individual landslides are 223 more difficult to delimit because they occurred in a 224 multiple and successive way, often one on top of an- 225 other. The main types in this area are rotational 226 rockslides in the upper part, often with multiple scarps 227 and debris slides in the lower parts, combined with 228 extensive rill and gully erosion. It is difficult to define 229 the age of the landslides, but the multiple and complex 230 landslide forms indicate that mass movements have 231 been active over a long period, associated with the 232 activity of the faults, and probably with relatively minor 233 individual displacements per event. The most recently 234 known landslide occurred in 1997, during an intensive 235 rainy season, but probably was caused by a leaking 236 water pipe; it destroyed a mini-hydroelectric power 237 plant. 238

Landslides in the denudational hills underlain by 239 metamorphic and terrigenous rocks (Fig. 3, A and B) are 240 generally smaller than those in the limestone hills, and are 241 not so related to tectonic lineaments. Most of them are 242 shallow rotational landslides occurring rather at random 243 on the steep hill-slopes or along river incisions. 244

3.2. Alluvial units and accumulational slopes 245

The Alluvial valley complex is related to the recent 246 sediments accumulated by to the principal river systems 247 (Figs. 3D and 2 for names). Sabanalamar river floodplain, 248 Macambo river floodplain and the most recent fluvial 249 channels in the Caujeri valley are part of this complex. 250 They are composed of alluvial and swampy deltaic 251 deposits where the rivers end in submerged valleys, due 252 to tectonic uplift. This complex is essentially a fluvial plain 253 with a combination of erosive and accumulation processes. 254 Accumulation prevails in the Sabanalamar and Macambo 255 floodplains and erosion in the Caujeri valley. Three 256 floodplain levels are recognizable: active, occasionally 257

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submerged and exceptionally submerged floodplains. The area is regularly inundated during intensive rain. Close to Sabanalamar Bay and in the surroundings of San Antonio del Sur town, there are brackish water lagoons and swamps. Mangrove vegetation is abundant only in the mouth of the Sabanalamar River.

The accumulational slopes are located in the northern 264 part of the coastal hills (Fig. 3). It is an intra-mountainous 265fluvio-marine deltaic plain approximately 2 km wide. 266Slopes slightly to the south (sea) side at 5 to 15° . The 267 materials range from colluvial, close to the mountains, to 268fluvio-marine. This complex seems to belong to an old 269 270planation surface, which collected all the sediments coming from the upward area, what is now Sierra de 271Mariana (Fig. 2), during the Pleistocene. The extension 272and volume of the Quaternary sediments reveals that 273 rainfall at that time was higher than currently. This might 274275be true considering the fact that in the northern border of the Sierra de Mariana the drainage seems to be cut-off due 276to large mass movements. In both the western and eastern 277sides of the area, the Pleistocene sediments are not 278 present. In the western side the drainage system was 279sufficiently strong to erode the sediments to the Baitiquirí 280 281 Bay, besides this part appears to be slightly more uplifted. In the eastern part the pre-Quaternary formation overlies 282 the ophiolites and the area has an irregular relief. The 283planation surface is raised 10 to 30 m above the current 284erosion levels, and around ten new channels have eroded 285the old plain and generated erosional scarps with the same 286height differences. The alluvial unit and accumulational 287slopes do not contain any major landslide features. 288

289 3.3. Coastal hills

These hills parallel the coastline (Fig. 3), with variable 290length and a width between one and 2 km. Three coastal 291 hills can be differentiated, between El Naranjo and 292Baitiquirí Bay, between the Baitiqurí bay and Sabanala-293mar Bay (Loma Los Aposentos) and between Sabanala-294mar Bay and Macambo town (see names in Fig. 2 and 295 relief in Fig. 3 F). The north side is totally covered by the 296Maquey formation (Phase1), mudstone and calcareous 297clay susceptible to landslides. The coastal slope, charac-298terized by marine terraces is composed of the Maya for-299 mation (Phase3) consisting of recent (Holocene) marine 300 deposits. These deposits act as "rings" of the coastal hills 301 and are uplifted between 5 and 10 m from current sea 302 level. 303

The different material and morphological characteristics on both sides of the coastal hills also result in different landslide types. Northern slopes are characterized by frequent, but small debris flows. On the coastal side, rockfalls occur in the marine terraces, caused by a 308 combination of karstic dissolution and physical weath-309 ering processes and triggered by wave erosion. Large 310 blocks with volumes between 15 to 40 m³ can be found 311 as part of the rockfalls. On top of the cliff various cracks 312 delimit the boundaries of future rockfall events. Magaz 313 et al. (1991) mapped three large rotational and two 314 translational rockslides in the marine terraces, covering 315 the entire seaward side of the coastal hills. Two of the 316 large rotational landslides are pre-Holocene because the 317 lower terrace was formed on top of the landslide toe. 318 The third landslide in Los Aposentos coastal hill 319 (Fig. 4A and B) is more recent than the others are 320 since the lower terrace (Holocene) was also destroyed. 321

3.4. Puriales de Caujerí depression

By far the most striking geomorphological feature in 323 the study area is the large oval shaped depression 324 (Puriales de Caujeri valley), which is considered to be a 325 graben with elevation differences up to 500 m. The 326 valley is limited on the west by a large scarp of the Sierra 327 de Caujeri, with some active retrogressive mass move- 328 ments. On the southern and northern parts the valley is 329 also surrounded by major fault scarps. The origin of the 330 Puriales de Caujeri depression can be interpreted as a 331 combination of tectonic and mass wasting processes. 332 The main fault systems (Sierra de Mariana and Sierra de 333 Caujeri) started to generate a graben depression after the 334 second Transgression-Regression period (Lower Mio- 335 cene to Late Miocene). After that, the 15-kilometer-long 336 N-S oriented Sierra de Caujeri carp, 300 to 400 m high, 337 has been the main area of landslide activity. In the north 338 the scarp ends 3 to 5 km north of Mameyal and has less 339 recent landslide activity. In the south the scarp intersects 340 the fault-controlled Sierra de Mariana scarp creating 341 another area with large landslides. Fig. 4 (C, D and E) 342 shows some of the landslides in the Caujerí scarp. 343

The most catastrophic landslide in the Sierra de 344 Caujerí scarp occurred after three days of heavy rain 345 during the passing of cyclone Flora on October 8, 1963, 346 the most devastating meteorological event known that 347 affected Cuba (see Fig. 4 D and E). A total of 1100 mm 348 of rainfall in three days was recorded in the Sierra de 349 Caujerí area (Trusov, 1989). The successive rotational 350 rockslide occurred in two pulses at about 45 min 351 intervals, which allowed some of the inhabitants to 352 escape, whereas 5–10 others were killed. No technical 353 report was made directly after the event although some 354 data were recorded during the fieldwork when a number 355 of interviews were held with some of the survivors. 356 However, due to the long time since the occurrence of 357

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Fig. 4. Different landslides from the study area. A (Spot Pan 28/12/1994) and B: Los Aposentos landslide, typical rotational landslide in the coastal hills; C (Spot Pan 28/12/1994), D (Aerial photo K10 survey, 02/02/1972), E: Jagueyes landslide, large landslide movements in the Sierra de Caujeri scarp.

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Fig. 5. Components of the heuristic landslide hazard model.

the landslide this information may no longer be reliable. 358 Most of the landslides in the Sierra de Caujerí scarp 359consist of a large scarp on the upper part, often up to 360 100 m high, which almost vertically cut the limestone 361 layer of the Yateras formation (Phase2), and the 362 underlying Maguey formation (Phase1). This scarp is 363 actually the back-scarp of multiple landslides, which 364 change from rockslide to debris flows. Around 150 365 different landslide events have been mapped along the 366 Sierra de Caujerí and Sierra de Mariana scarp (Figs. 2, 3 367 368 and 4).

369 4. Qualitative assessment of landslide hazard

The geomorphological mapping provided in-depth 370 knowledge of the causal factors for landslides in the study 371 area and was used to assess landslide susceptibility. 372 Qualitative weighting, one of the heuristic methods 373 (Soeters and Van Westen, 1996), was selected, given the 374relative small scale, the available data, and the character-375 istics of the study area. Besides, the TMU mapping may 376 produce biased results when using a statistical method due 377 to the high spatial correlation between the landslides 378 inventory and some units in the TMU map. 379

The analysis was carried out in different steps fol- 380 lowing the decision-support system method, according 381 to the analytic hierarchy process (Saaty, 1980; Saaty and 382 Vargas, 2001). First the various components (causal 383 factors) in the model were selected, and their hierarchi- 384 cal relation determined. Weights then were assigned to 385 these maps, in a standardized way. A combination for- 386 mula integrated the weights to produce a final map, 387 which was parsed into a number of classes (Bonham- 388 Carter, 1996). 389

The design of a heuristic landslide hazard model 390 requires thorough analysis of the causative factors and 391 their relationship in the study area. That was done based on 392 the geological and geomorphological interpretation along 393 with the fieldwork campaign. The selection of the model 394 components was organized in a tree-shaped structure 395 (Fig. 5). The upper most general level of the components is 396 called criteria, and consists (in ranked order) of Geomor-397 phology, Topography, Geology, Tectonics, and Hydrolo-398 gy. These criteria were further subdivided into nine 399 variables, specific attributes, such as slope, internal relief 400 and slope shape for Topography. The variables are 401 described in Table 1. The relation of each to landslide 402 occurrence can either be favorable or unfavorable.

t1.1 Table 1

t1.2	Variables for the heur	istic model			
t1.3	Variable	Origin	Scale	Units	Relation
t1.4	Slope	From the original DEM using the method of ILWIS software package (ITC, 2001)	Interval	Degrees	Favorable
t1.5	Internal relief	From the original DEM using the method of ILWIS software package (ITC, 2001)	Ratio	Meters/	Favorable
				hectares	
t1.6	Shape	From the original DEM using the method of ILWIS software package (ITC, 2001)	Ratio	N/A	Favorable
t1.7	Geology	By reclassifying the TMU map	Categorical	N/A	N/A
t1.8	Faults	Calculating a distance from the fault map and dividing into four classes	Ratio	Meters	Unfavorable
t1.9	Springs	Calculating a distance from spring points and dividing into four classes	Ratio	Meters	Favorable
t1.10	Drainage distance	Calculating a distance from the drainage map and dividing into four classes	Ratio	Meters	Favorable
t1.11	Geomorphological	By reclassifying the TMU map	Categorical	N/A	N/A
	subunits				
t1.12	Landslides subzones	By reclassifying the TMU map	Categorical	N/A	N/A

t1.13 See text for explanation (N/A — not applicable).

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The variables selected needed to be standardized in 404 order to properly compare them. The values of the 405 variables were converted into a number of classes, 406 depending on the type of variable. Every class received 407 a value between 0 and 100, considering that the value of 408 all classes in one variable need to sum up to 100. Class 409 boundaries for numerical variables were selected from the 410 25-cumulative percentage intervals of their histogram. 411 The classes of the categorical variables (e.g. geology) 412 were "weighted" according to a ranking method (see 413 below). 414

Once all the variables were standardized, weights 415 416 were assigned to the corresponding levels of criteria and variables in three different ways: directly by expert 417 opinion, by pairwise comparison matrix and by ranking. 418 The weight values range between 0 and 1 and need to 419 sum up to 1 among the variables within a criterion and 420 among the criteria. For checking the weight assignment 421 the decision-support system called DEFINITE was used 422 (Janssen and Van Herwijnen, 1994). In the first method 423 the weights of the criteria and variables were assigned 424 directly based on expert opinion and field experience. 425For the pairwise comparison matrix, each variable (or 426 427 criterion) is compared to all others in pairs in order to evaluate whether they are equally significant, or whether 428 one of them is somewhat more significant/better than the 429other for the goal concerned. In the ranking method the 430 criteria and variables are simply ranked according to 431 their importance as landslide controlling factors. The 432 rankings can be considered units on an ordinal scale. 433 Consequently the weights can be found by standardizing 434 the rank order (Voogd, 1983). The values for criteria, 435variable and classes were tabulated using Microsoft 436

t2.1 Table 2

2.2 Weight for criteria and variables for three method
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Components	Direct method		Pairwise matrix		Ranking method	
Topography	0.3		0.224		0.257	
Slope		0.7)	0.7		0.7
Internal relief		0.2		0.2		0.2
Shape		0.1		0.1		0.
Geology	0.2		0.131		0.157	
Formation		1		1		1
Tectonic	0.05		0.040		0.065	
Active faults		1		1		1
Hydrology	0.05		0.038		0.065	
Springs		0.5		0.5		0.
Drainage density		0.5		0.5		0.
Geomorphology	0.4		0.566		0.457	
Subunits		0.4		0.4		0.4
Landslides zones		0.6		0.6		0.
Total for criteria	1		0.999		1.001	

Excel and a simple summation formula applied to 437 interactively evaluate the effect of the weights on the 438 overall weight of the qualitative landslide hazard.

The three weighting methods gave comparable results, 440 as can be seen from Table 2. For the pairwise comparison 441 matrix method the inconsistency value was 0.08, 442 demonstrating that the weights are sufficiently reliable. 443 The inconsistency parameter measures randomness of the 444 expert judgments, and ranges from 0 to 1. As a conclusion 445 the initial weights assigned by expert opinion were taken 446 for the analysis. 447

To calculate a final weight for a single pixel the weight 448 of the three components (class value, variable and 449 criterion) was simply multiplied. For example, the slope 450 class "shp4" (>20.4°) has a weight of 50, which was 451 multiplied by the weight of the "Slope" variable (0.7) and 452by the weight of the "Topography" criterion (0.3), so that 453 the final weight was 10.5. Final weights were calculated 454 for each variable and resulted in separate layers, which are 455 generated with the CARIS GIS software package 456 (Universal Systems Ltd, 2000) for the spatial overlay 457 analysis. In this GIS, each class is created in raster format 458 by different options according to the map characteristics: 459 by buffering (e.g. distance to faults), by selecting 460 polygons (e.g. TMU subunits) and by reclassifying 461 pixel values in a raster image (e.g. slope angle values). 462 All weight maps were added into a layer with final 463 weights. 464

The final weights of the resulting map ranged from 0.5 465 to 47.1. This map was classified into 10 divisions 466 interactively, during which the relation with existing 467 landslide areas and geomorphological units was evaluat- 468 ed. Although the map gives a good indication of the 469 qualitative landslide hazard in the study area, too many 470 classes might make it difficult to use by decision makers 471 for development planning. Therefore, the hazard map has 472 ten classes, which are grouped into three simplified 473 categories: high, moderate and low (Fig. 6 and Table 3). 474

The final hazard map was combined with the TMU 475 map-database in order to obtain information about the 476 attributes for each hazard class. Table 3 shows the leg- 477 end of the final landslide hazard map with statistical 478 information related to the number of landslides and 479 landslide density per class. 480

Additionally, in the final hazard map the flooding 481 areas were included with two categories: flooding zones 482 (F1) and dam break zones (F2) (Fig. 6 and Table 3). 483 Flooding zones were primarily acquired from a disaster 484 management plan used by the local civil defense 485 authority. The flooding boundaries were corrected by 486 photo-interpretation, DEM analysis and fieldwork. Dam 487 breaks zones were obtained from the dam project reports 488





Elements at risk

	House, school or cemeter
_	National Road
_	Provincial Road
_	Paved Road
	Unpaved Road
	Trail
	Coast line

Fig. 6. Landslide hazard map. San Antonio del Sur, Guantánamo, Cuba. See Table 3 for explanation of the legend.

made by the Institute of Water Resources for Guatánamo
province. The two areas mapped are related to Pozo
Azul and Palmarito dams. Dam location and names
appear in Fig. 2. More detail information about the
flooding zones was available, but could not be displayed
at 1:50,000 scale.

5. Qualitative assessment of landslide risk

Too little information was available to carry out a more 496 sophisticated landslide risk assessment. In particular, 497 there were not enough data on the probability of landslides 498 of different magnitudes to make a (semi) quantitative risk 499

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t3.1 Table 3 t3.2 Characterization of the 10 landslide hazard classes and 2 flooding hazard classes (see Fig

Hazard	Overall class	Hazard class	Weight range	Area (ha)	Number of TMU units	Number of landslides	Area of landslides (ha)	Landslide density (%)	General hazard description
Landslides	Low	H1	0.50- 5.16	8669	88	0	0	0	No landslides expected. Areas can be corridors for mudflows or other intensive
		H2	5.17- 9.82	4176	40	0	0	0	mass wasting processes. In some parts small landslides can
		H3	9.83– 14.48	18,107	125	8	210	1.1	happen in extreme conditions. The areas are suitable for development projects.
	Moderate	H4	14.49– 19.14	18,361	94	22	338	1.8	Moderate to high possibility of landslides occurrence during intensive or prolonged
		H5	19.15– 23.80	2013	85	46	901	44.7	rainfall. These areas contain most of the existing landslide zones. Most of the landslide
		H6	23.81– 28.46	1,702	87	79	1,428	83.8	materials are unconsolidated and susceptible to being reactivated in smaller proportions. More studies are required for development of the area, Land use changes should be previously studied in relation to landslide hazard problem.
	High	H7	28.47- 33.12	969	77	76	963	99.3	High to very high landslide hazard areas. A high possibility of landslide occurrence
		H8	33.13– 37.78	719	59	59	715	99.5	during rainy conditions. No development is recommended in these areas.
		H9	37.79– 42.44	291	36	36	289	99.4	Possible relocation of land for agricultural use. Highly recommended re-allocation
		H10	42.45– 47.10	439	24	24	438	99.8	of existing population in these areas.
Floods		F1	N/A	1325	15		0	0	Flooding areas up to 5 years return period taken from local civil defense authority and updated yearly. Appropriate warning system needs to be maintained. The area could be used for seasonal agricultural products. Land-use planning should consider flooding hazard limits to re-allocated existing infrastructure and avoid new developments.
		F2	N/A	503	6	0	0	0	Dam break flood limit taken from dam project report. Engineering conditions of the dam should continuously be checked. The area could be used for agricultural products. Land-use planning should consider this hazar limit to re-allocated existing infrastructure an avoid new developments.

assessment. Therefore only a basic qualitative analysis 500was carried out through the combination of the landslide 501hazard map with basic elements at risk. During the 502fieldwork, information was collected on buildings and 503roads. As the study area is a rural environment, most of the 504buildings are isolated farmhouses; a number of small 505schools and medical centers also were identified. Most 506roads are unpaved country roads, except for the main road 507 in the south of the study area, which runs along the coast, 508mostly on the land-side of the coastal hills. The roads and 509 buildings are indicated in Fig. 6. No attempt was made to 510evaluate the potential losses of crops. In the study area 511there are a few small settlements, of which San Antonio 512

del Sur is the largest. Other villages and hamlets are 513 Puriales de Caujerí, Baitiquirí, Macambo, Guaibanó, and 514 El Naranjo (see names in Fig. 2). 515

MOST of the 3317 buildings in the study area are made 516 of wood over a concrete base, and few are of masonry only 517 or of reinforced concrete cement (RCC), which are nearly 518 all found in the settlements. There are 88 small schools 519 scattered over the area, 6 medical centers and 4 small 520 cemeteries. Building density is highest in the central-north 521 part of the Caujerí depression, where many buildings are at 522 risk in case a large landslide originates from the Caujerí 523 scarp. To analyze the qualitative risk level of the buildings, 524 GIS overlay was calculated between the building map and 525

the hazard map. The result is shown in Table 4. Most 526buildings and all of the small schools are located in low 527and moderate landslide hazard zones. No buildings are 528 currently in the highest hazard zones. A similar analysis 529was made for the road network (see Table 4). The 530provincial and national roads do not cross the high hazard 531 areas, although quite a few roads are in the moderate 532hazard classes. To refine these results in the future more 533data related to the elements at risk should be collected, 534including population distribution, and the replacement 535 value of buildings and roads, as well as the value of 536different agricultural crops. 537

538 6. Discussion and conclusions

Detailed geomorphological mapping provides infor-539 mation on the site-specific conditions under which 540different landslide types occur in different parts of the 541study area. Landslides are concentrated along the Caujerí 542scarp, but also in the coastal hills and in the northern part 543of the Baitiquirí area. There, the landslides have different 544characteristics and causative factors. Subdivision of the 545 terrain into 603 terrain mapping units, individual 546 547 homogeneous polygons, allowed for a more detailed characterization of the terrain than would be possible 548 through conventional map-overlay of main factor maps. 549The boundaries of different landslide parts were surveyed 550by photo-interpretation, and landslides were described in 551 the field from a detailed checklist, where information was 552collected related to landslide type, subtype, relative age, 553and depth. 554

This data set enable us to generate a heuristic model, using multicriteria analysis, which was successful in classifying the area into different hazard classes, which can be displayed either in 10 more detailed or 3 generalized classes, depending on the user's map needs. Improvements 559 in the method could be made if different weight maps were 560 produced for different landslide types. However, since the 561 geomorphological units are one of the main variables in 562 the heuristic analysis, this could be partly taken into 563 account, by assigning weights to each individual unit. 564

Overall landslide qualitative hazard can be considered 565 rather low, with the exception of a number of specific 566 areas, such as the Caujeri scarp, and the area directly 567 below. However, due to the absence of a sufficiently long 568 landslide record, it was not possible to indicate probabil- 569 ities assigned to the 10 classes. Historical landslide 570 information is not available, except for a few isolated 571 landslides. It is known that large landslide can happen 572 along the Sierra de Caujeri, similar to the failure that 573 occurred in 1963 during the passage of a hurricane. The 574 situation in the study area can be considered as 575 representative for the whole of Cuba, where due to the 576 lack of a landslide inventory, the knowledge about 577 landslide causative factors and mechanisms is still limited. 578

Recently, however, the National Civil Defense and 579 the Ministry of Science, Technology and Environment 580 have decided to establish a system for landslide risk 581 assessment in the Cuban Archipelago. The system will 582 include the design and implementation of a national 583 landslide inventory database and landslide risk-assess- 584 ment procedures at different disaster management levels 585 (Castellanos and Van Westen, 2005). 586

An important component of this system will be the 587 involvement of local staff of the Civil Defense at the 169 588 municipal centers, including San Antonio del Sur. A 589 simple landslide reporting form has been designed, and 590 workshops will be conducted to train the staff and make 591 them aware of the procedure. Once the local officers report 592 a landslide, a landslide expert from the central office will 593

t4.1 Table 4

t4.2 The number of buildings and length of roads per hazard class

t4.3	Hazard classes	Buildings (nr)	Houses (nr)	Schools (nr)	Cemeteries	National roads (km)	Provincial roads (km)	Paved roads (km)	Unpaved roads (km)
t4.4	H1	1370	1338	29	3	33.3	27.9	29.4	151.4
t4.5	H2	381	367	13	1	2.4	0	17.6	48.9
t4.6	H3	588	570	18	0	5.6	20.7	15.7	132.4
t4.7	H4	947	919	28	0	0.1	1.2	7.4	120.6
t4.8	H5	69	68	0	1	0.1	0	0	16.6
t4.9	H6	36	36	0	0	0.6	0	0.9	5.8
t4.10	H7	30	29	1	0	0	0	0	7.9
t4.11	H8	14	14	0	0	0	0	0.7	0.8
t4.12	H9	0	0	0	0	0	0	0	1.1
t4.13	H10	1	1	0	0	0	0	0	0
t4.14	Total	3436	3342	89	5	42.1	49.8	71.6	484.6
t4.15	Actual nr.	3409	3317	88	4				

t4.16 The small difference between the total and the actual number is due to processing errors in the rasterization procedure.

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visit the site and complete the questionnaire in more detail. 594Such a system for landslide data collection might be less 595effective in other countries, due to insufficient reporting 596staff at the local level. In Cuba, however, the Civil Defense 597is well organized and very effective as reflected in a 598comparison of disaster-related casualties numbers in Cuba 599with those of neighboring countries such as Haiti or the 600 Dominican Republic (Thompson and Gaviria, 2004). 601

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